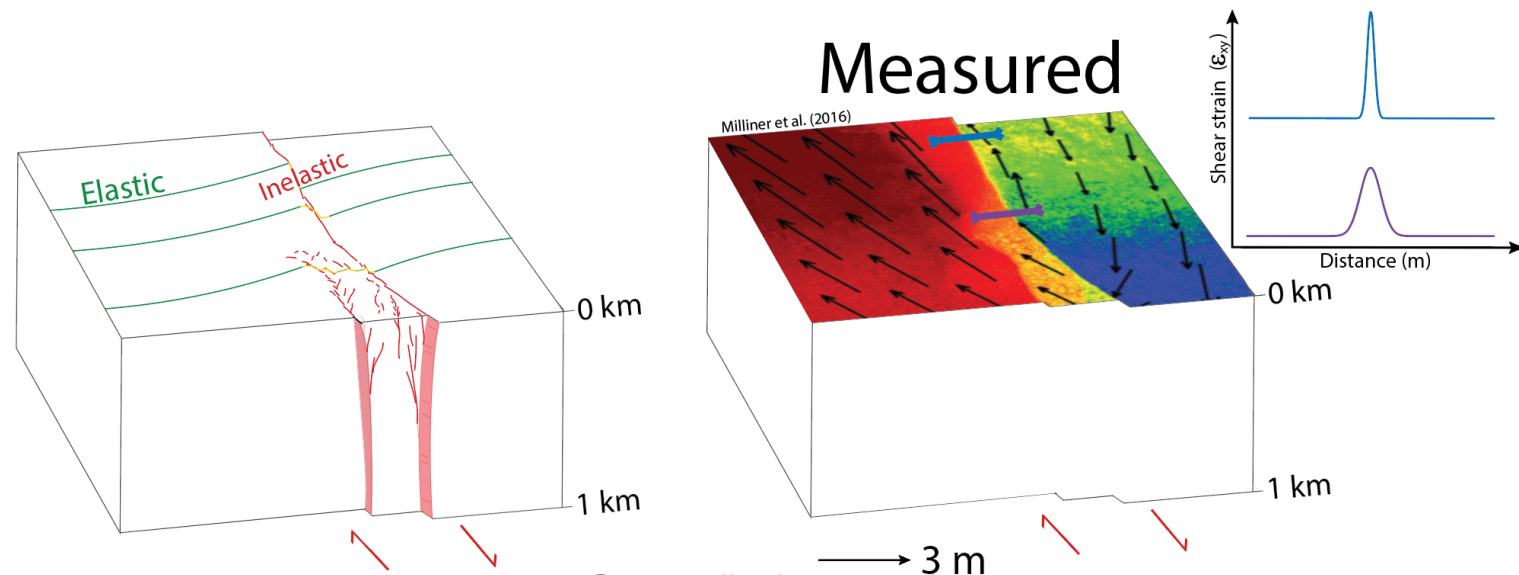


High Resolution Geodetic Measurements of Co-seismic Fault-zone Deformation for PFDHA

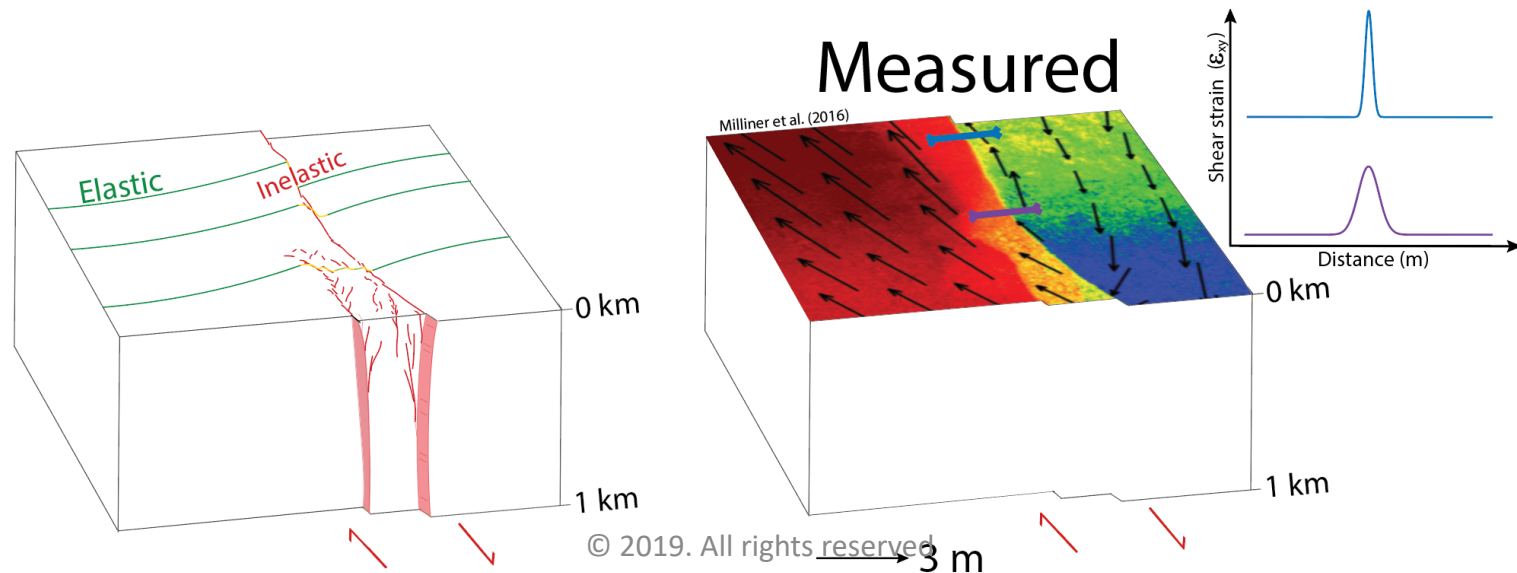
Chris Milliner, Postdoc, Jet Propulsion Laboratory, California Institute of Technology

Rui Chen, California Geological Survey



Introduction

- **Aim:** Use high-resolution geodetic data that can resolve near-field surface deformation to improve PFDHA models.
- **Motivation:** Measuring distributed faulting is highly challenging in the field (largely due to lack of cultural features that span the fault zone in perpendicular manner).
- **Method:** Use optical and SAR pixel tracking from multiple $M_w > 7$ surface rupturing earthquakes to measure distribution of strain across fault.



Multiple earthquakes

Event	Year	Mw	Length (km)	# strain profiles	Mechanism
Landers	1992	7.3	80	~1000	SS
Hector Mine	1999	7.1	50	~700	SS
EMC	2010	7.2	120	500-1500	SS, normal
Balochistan	2013	7.7	240	30	SS
Napa	2014	6.1	30	30-150	SS
Kumamoto	2016	7.1	40	40	SS, normal
Kaikoura	2017	7.8	120	120-600	SS, thrust
Canterbury	2011	Mw 6.2			SS,
Norcia, Amatrice					SS, normal
China x ??? Gareth F.?					??
Papau New Guinea					thrust
Palu	2018	7.5	150	70-500	SS

Methods: Optical & SAR pixel tracking

Correlate optical radiometric data (visible EM)

COSI-Corr: Co-registration of Optically Sensed Images and Correlation



1 Satellite imagery
acquired at different
times, any resolution,
possibly by different
sensors

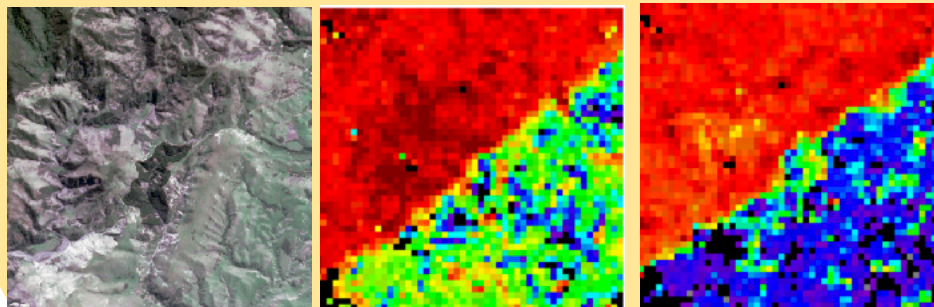
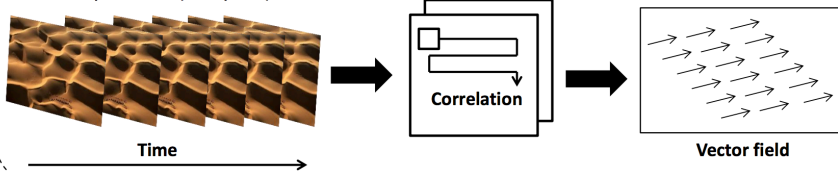
Measuring surface
processes
from optical imagery

COSI-Corr

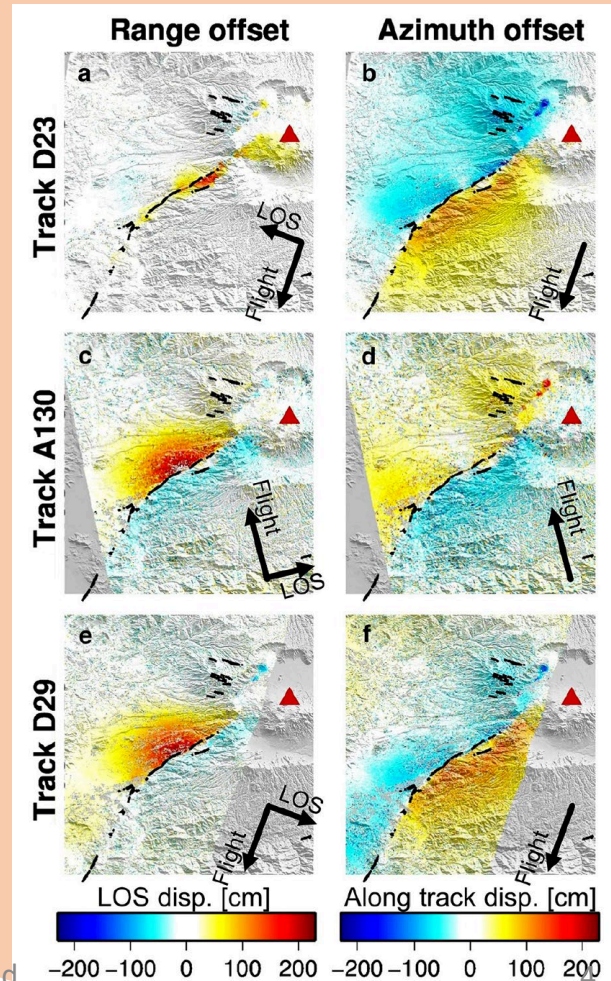
2 Automatic registration
with accuracy of 1/10 of
the pixel size (sub-pixel)

3 Automatic comparison of
images to measure motion

4 Ground
deformation

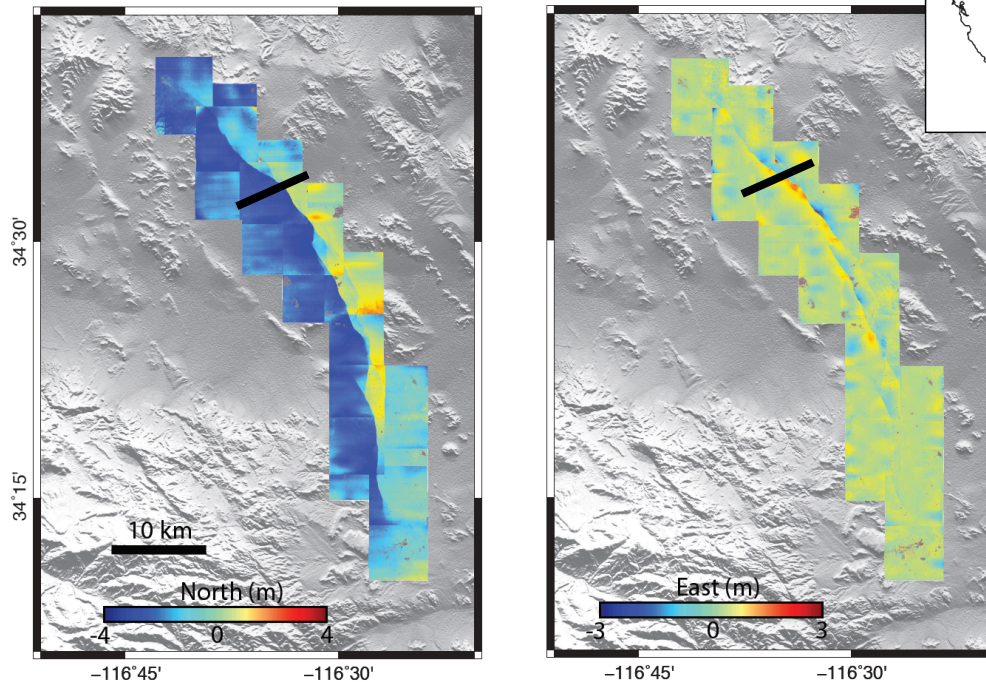


Correlate amplitude of radar
backscatter (microwave EM) > 3
look directions → 3D motion



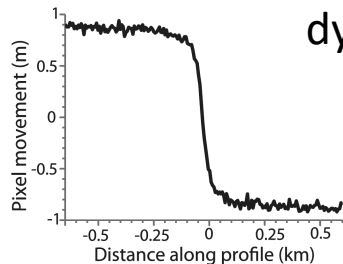
Results – Optical pixel tracking – 2D

1992 Mw 7.3 Landers

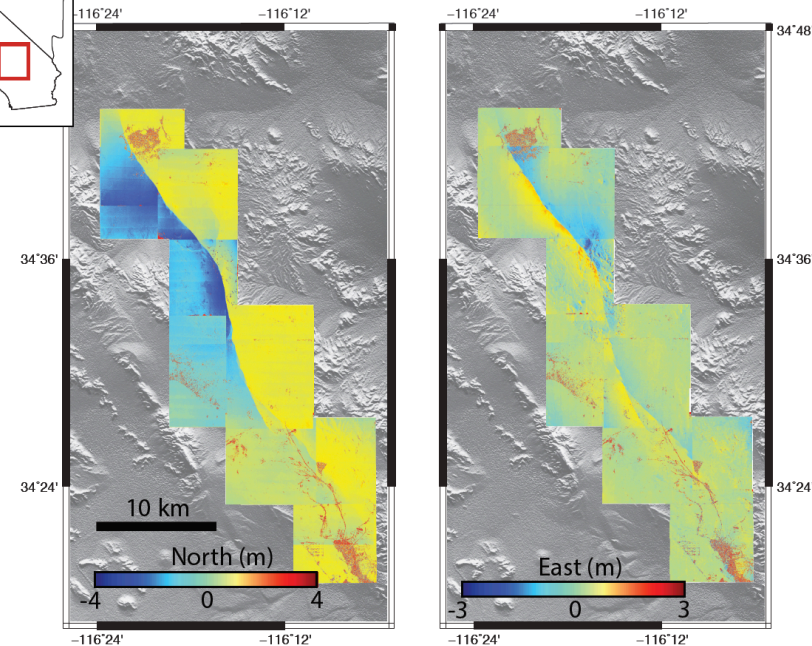


dx

dy



1999 Mw 7.1 Hector Mine

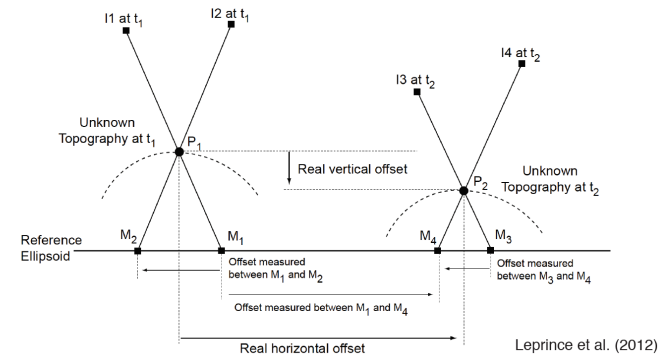


dx

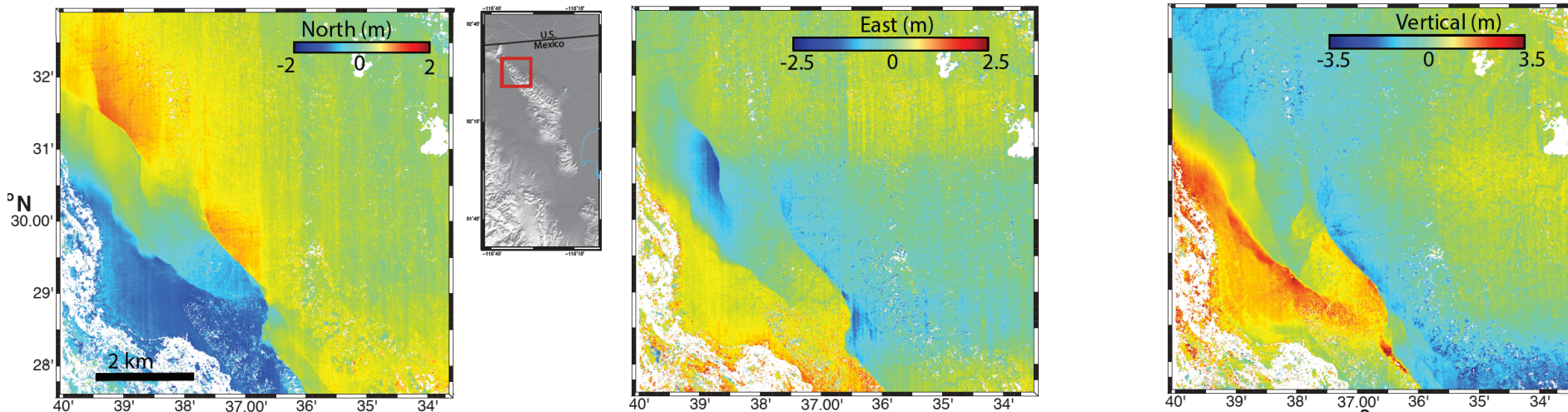
dy

Results – Optical pixel tracking – 3D

- El-mayor Cucapah, M_w 7.2, 2010, Mexico
- Oblique: strike-slip, normal
- Rupture length: 120 km



2010 Mw 7.2 El-Mayor Cucapah



Leprince et al. (2015)

dx

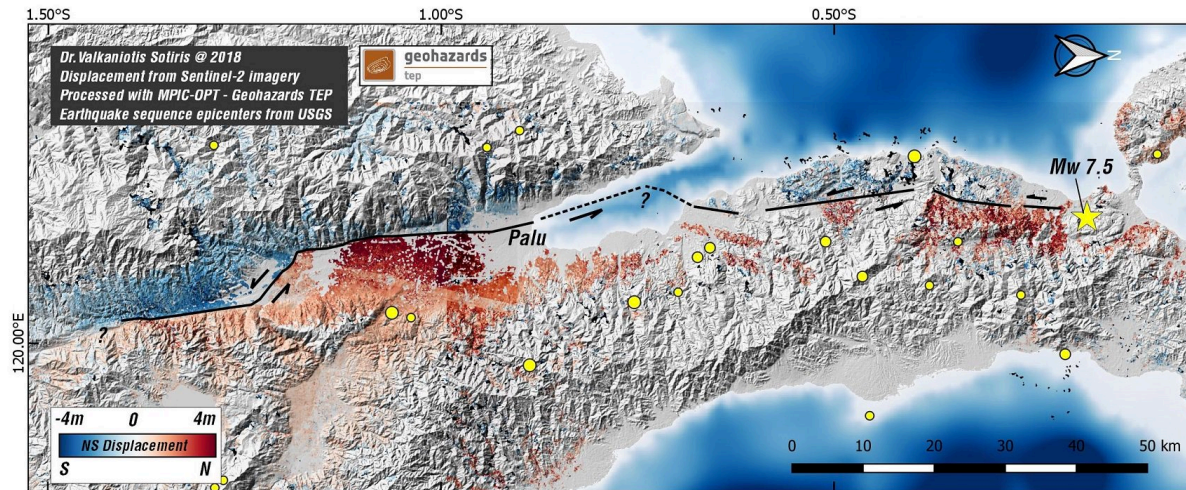
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dy

dz

Palu, Indonesia, Mw 7.5

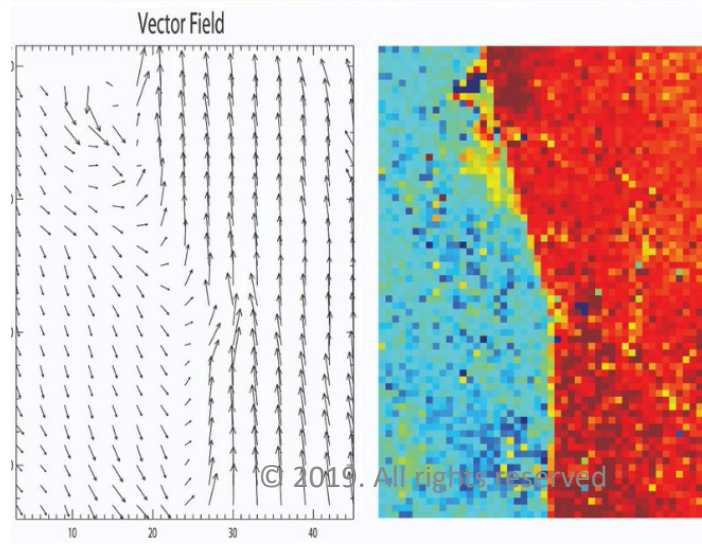
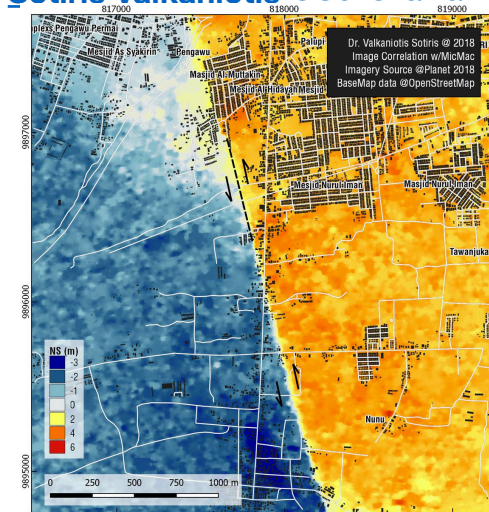
- 150 km surface rupture
- Sentinel 2 data



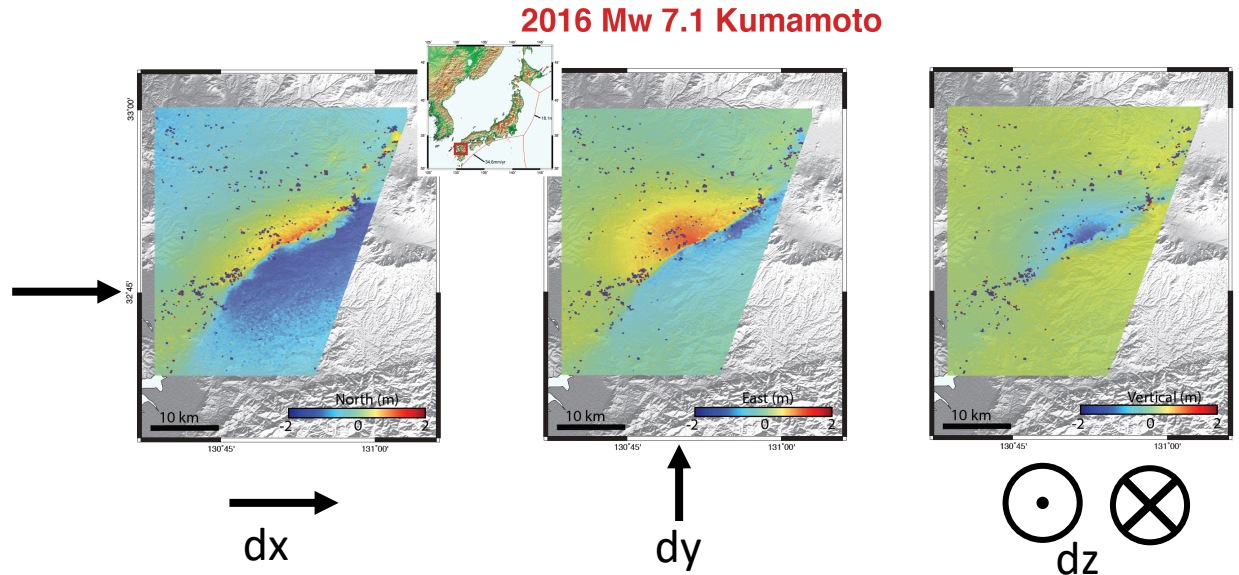
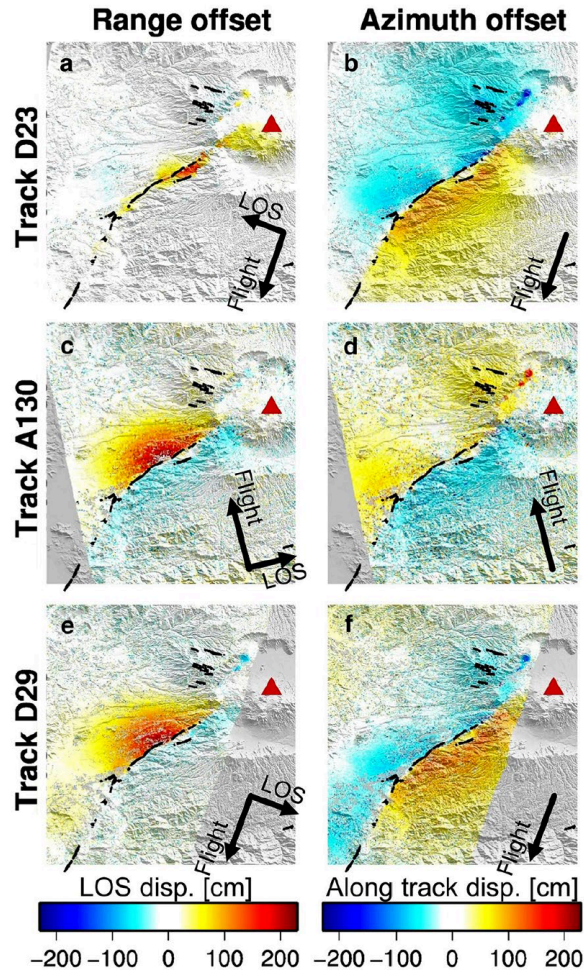
Sentinel 2 – 10 m resolution



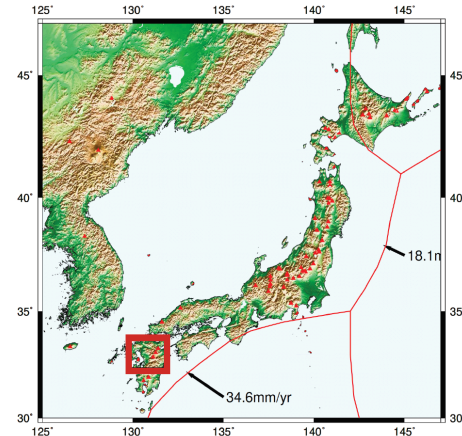
Sotiris Valkaniotis @SotisValkan



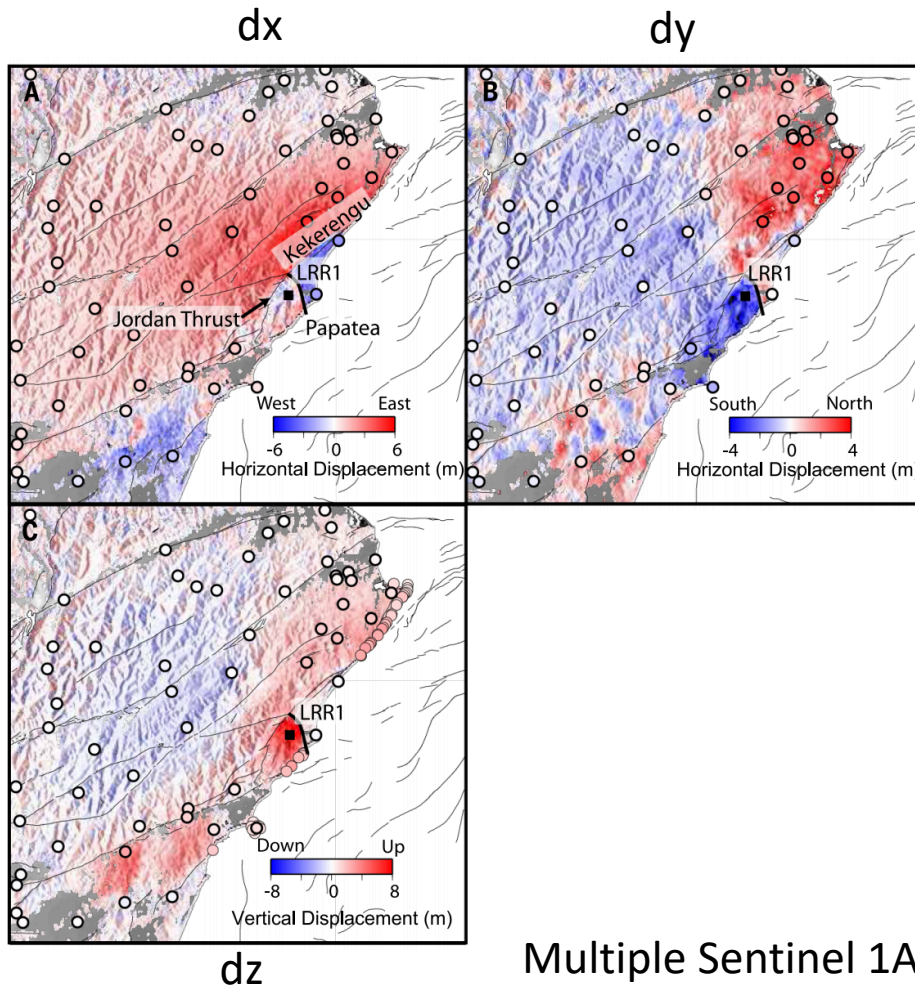
Radar pixel offsets – 3D



- Japan, 2016, Mw 7.1
- Oblique: Strike-slip, normal
- Rupture length: 40 km
- Noise level ~ 20 cm (1 sigma)
- Pixel resolution: ~ 25 m



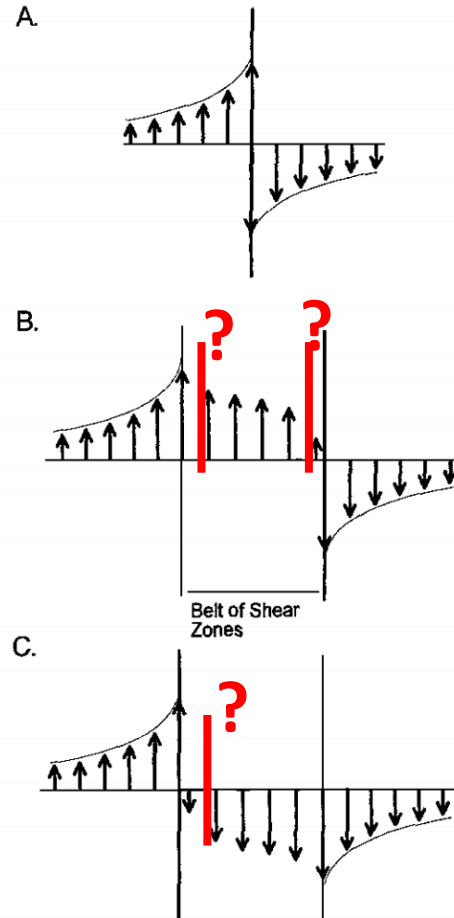
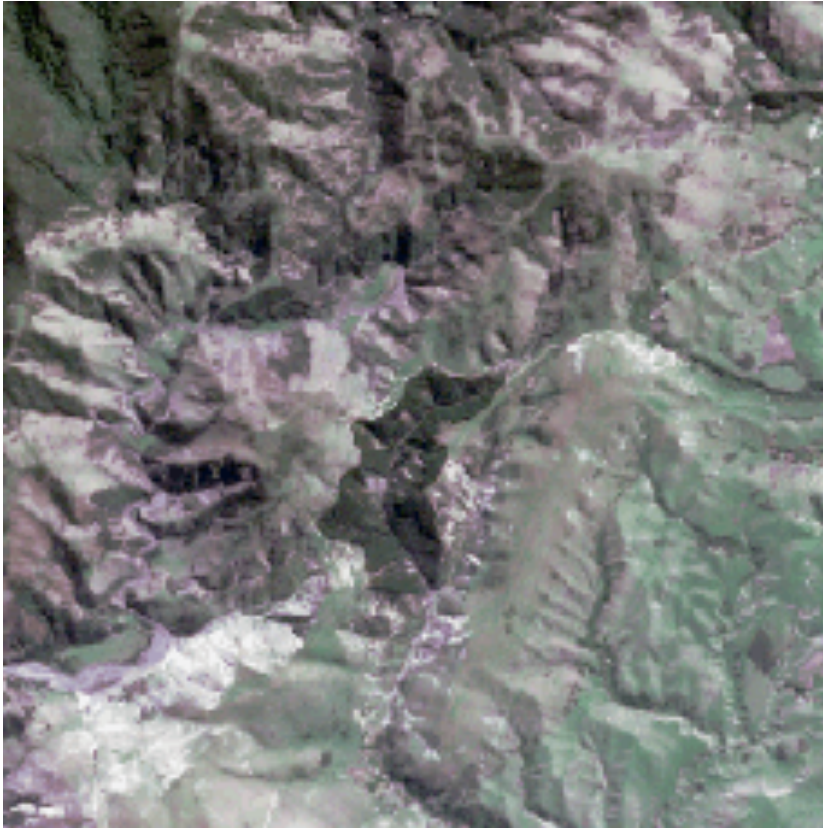
2017 Mw 7.8 Kaikoura, NZ (?) – 3D SAR



- Highly complex event, > 12 major faults.
- Underlying mega-thrust thought to participate, and perhaps primarily control rupture propagation and explain the large rupture complexity.
- Due to uniqueness of rupture, debate whether to include this in PFDHA?

Multiple Sentinel 1A (ESA) radar scenes,
C-band (3 cm wavelength)

How to calc. probabilities using geodetic data?



Johnson et al. (1994)

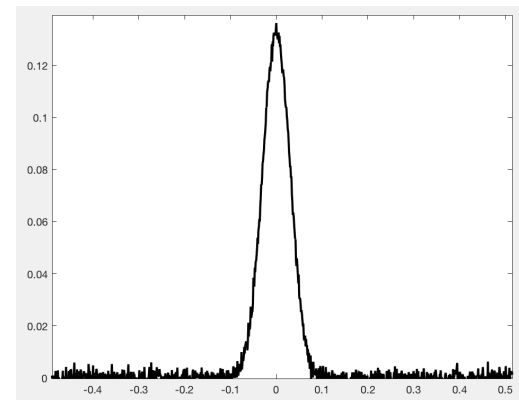
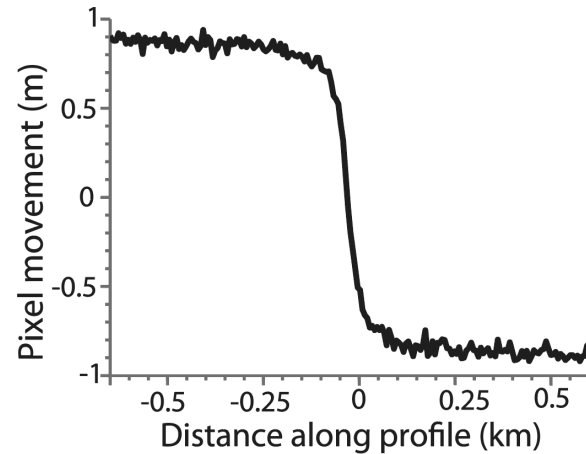
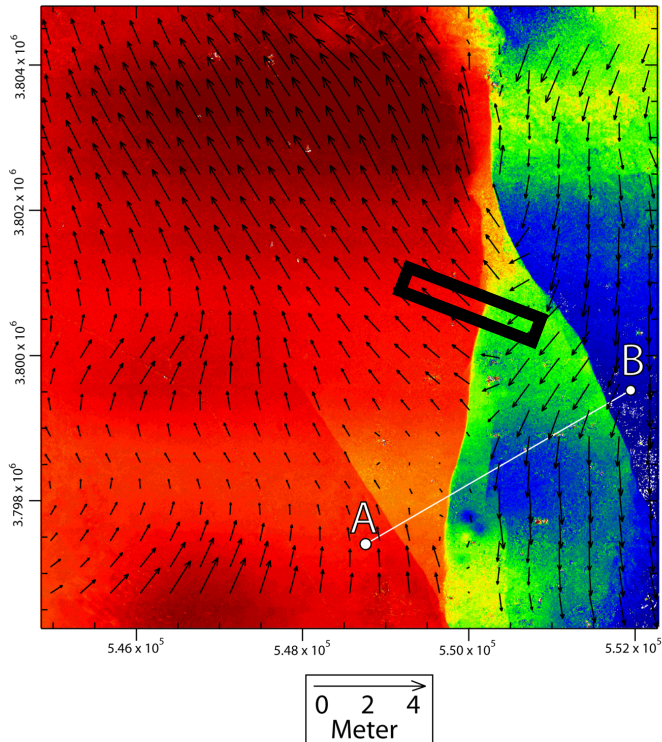
Key assumptions:

- We use strain, not displacement on indiv. fractures.
- Due to geodetic imagery averaging spectral properties over an area + corr. window → cant resolve individual fractures, the velocity field is almost continuous.
- Therefore product we'll provide is the amount of shear strain a structure will experience over a given length scale == total displacement.

$$\lambda(\varepsilon \geq \varepsilon_o)_{xyz} = \alpha(m) P[sr \neq 0 | m] \int_r P[\varepsilon > \varepsilon_{inelastic} | r, z] P[\varepsilon \geq \varepsilon_o | r, \varepsilon_{inelastic}] dr$$

Results - profiles

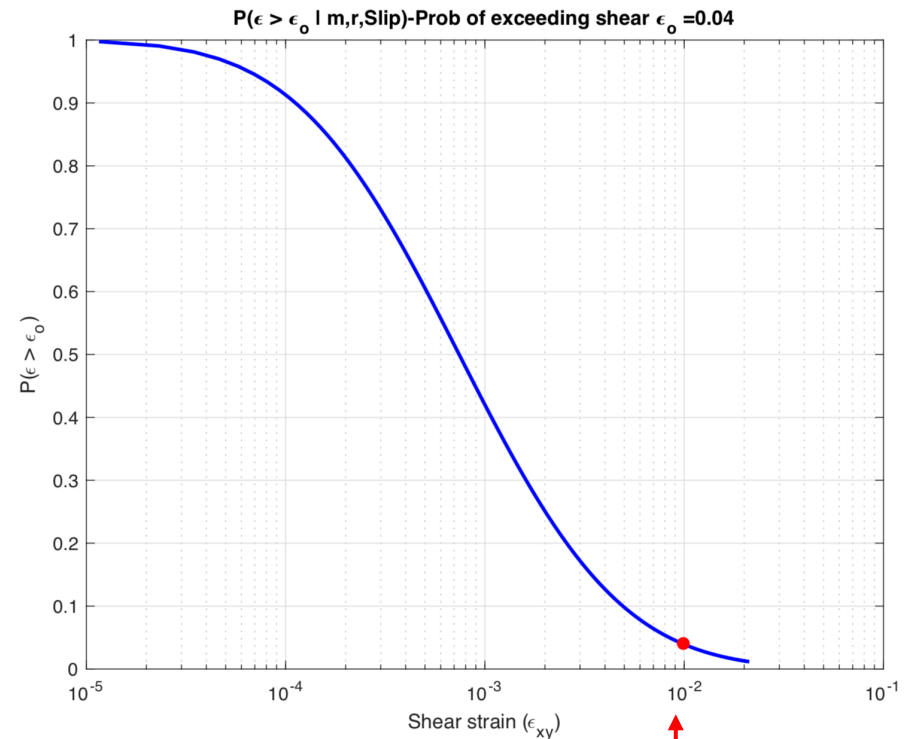
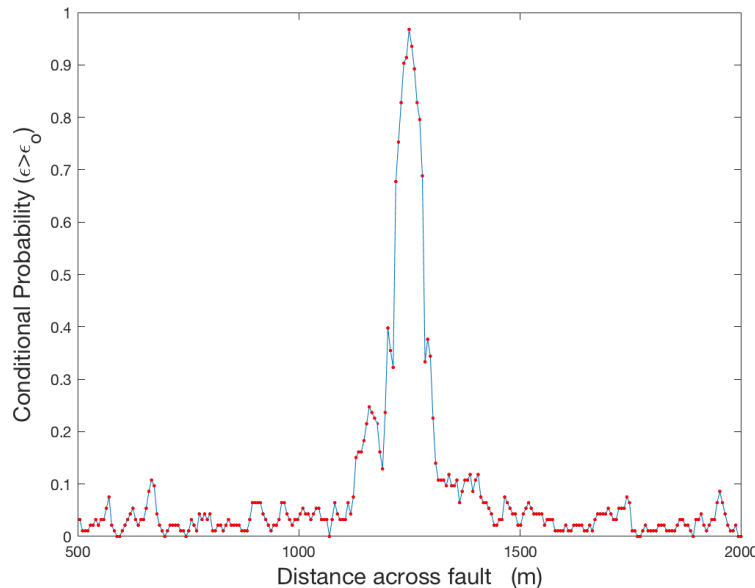
Vector Field Showing Ground Deformation at the Kickapoo Stepover



1. Draw stacked profiles perpendicular to fault (~200 m width) → fault parallel motion
2. Calculate gradient → shear strain

Results - Probability calculation

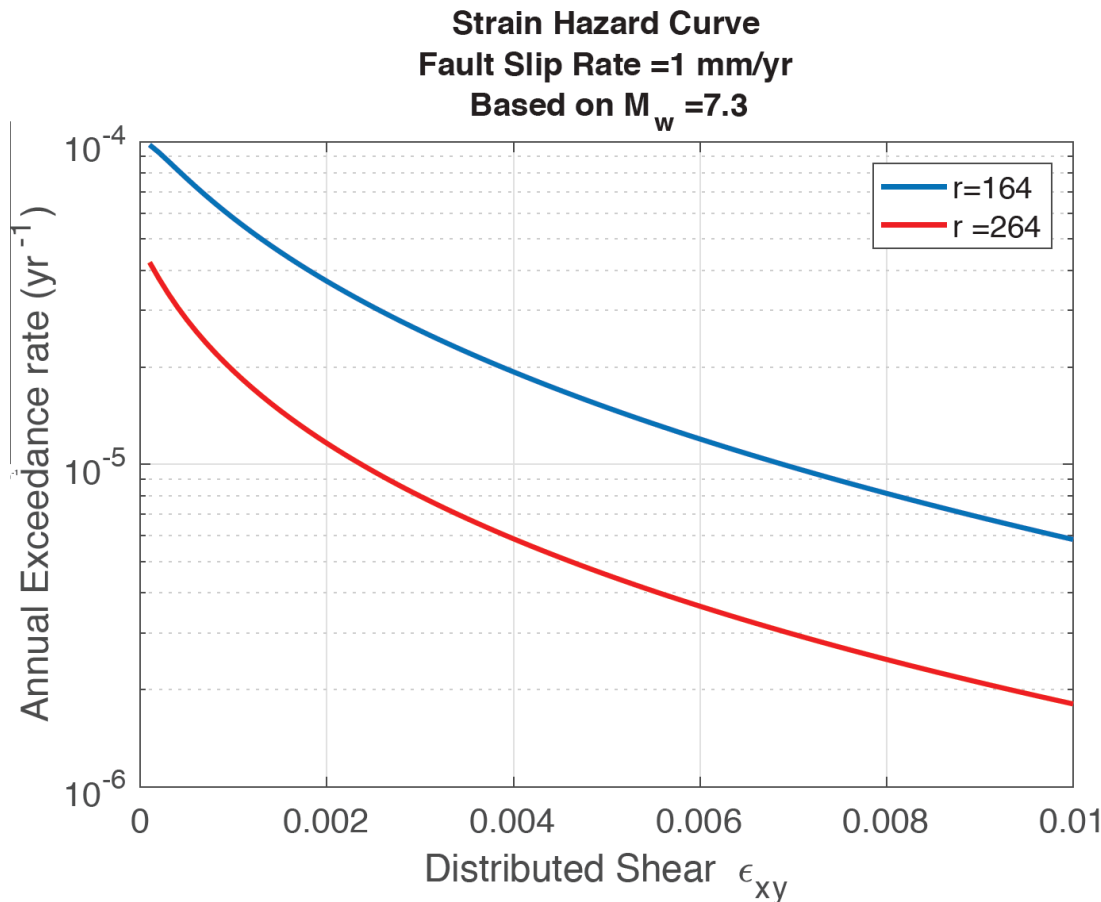
$4\text{e-}3$ = yield strength of granite \rightarrow
conservative inelastic strain



$$\lambda(\varepsilon \geq \varepsilon_o)_{xyz} = \alpha(m) P[sr \neq 0 | m] \int_r \underbrace{P[\varepsilon > \varepsilon_{inelastic} | r, z]}_{\text{blue curve}} \underbrace{P[\varepsilon \geq \varepsilon_o | r, \varepsilon_{inelastic}]}_{\text{red dot}} dr$$

Similar to Petersen et al. (2011),
but using strain

Results – Hazard curve



Scenario event:

- Assuming fault slip rate = 1 mm/yr
- Magnitude = 7.3
- Hazard of strain for 2 distances from main rupture.
- Another key assumption:
 - Location of primary rupture has been identified (with confidence from trenching) uncertainty on location not considered

Conclusions

- **Aim:** outline a standard method for high-res geodetic data to constrain PFDHA models.
- **Results:** Geodetic data allows us to gather thousands of strain profiles.
 - Potential to do PFDHA for SS, normal + thrust
- **Assumptions& Limitations:**
 - We quantify strain, not displacement on individual fractures.
 - Can't discern elastic vs inelastic, we have to assume a threshold value that exceeds yield strength, or can let user decide the minimum strain to exceed.
 - Data of varying resolution + noise → varying sensitivities to strain.

Future work

- **Going forward:**

1. Most data already gathered
2. More eq's (1-4)
3. Include Kaikoura?
4. Separate oblique faulting events?
 1. Decreases number of data per faulting style
5. Assess whether near-surface geology, fault geometry, sediment thickness etc... has an effect → this could reduce epistemic uncertainty.

- **What we need (data):**

- 2018 M_w 7.5 Palu, Indonesia – Planet labs (free)
- 2013 M_w 7.7, Balochistan, Pakistan – Landsat (free)
- 2014 M_w 6.1 Napa, US – lidar, optical (pre-existing)

- **Timeline:**

1. Will verify PFDHA code with Rui Chen - visit Sacramento next month.
2. Process more data (<2 months).
3. Publish, < 1 yr timeframe - a method detailing how to use geodetic data for PFDHA + present results from multiple earthquakes.